



21st century hydrological modelling for optimizing ancient water harvesting techniques

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water harvesting techniques: often go back to ancient times





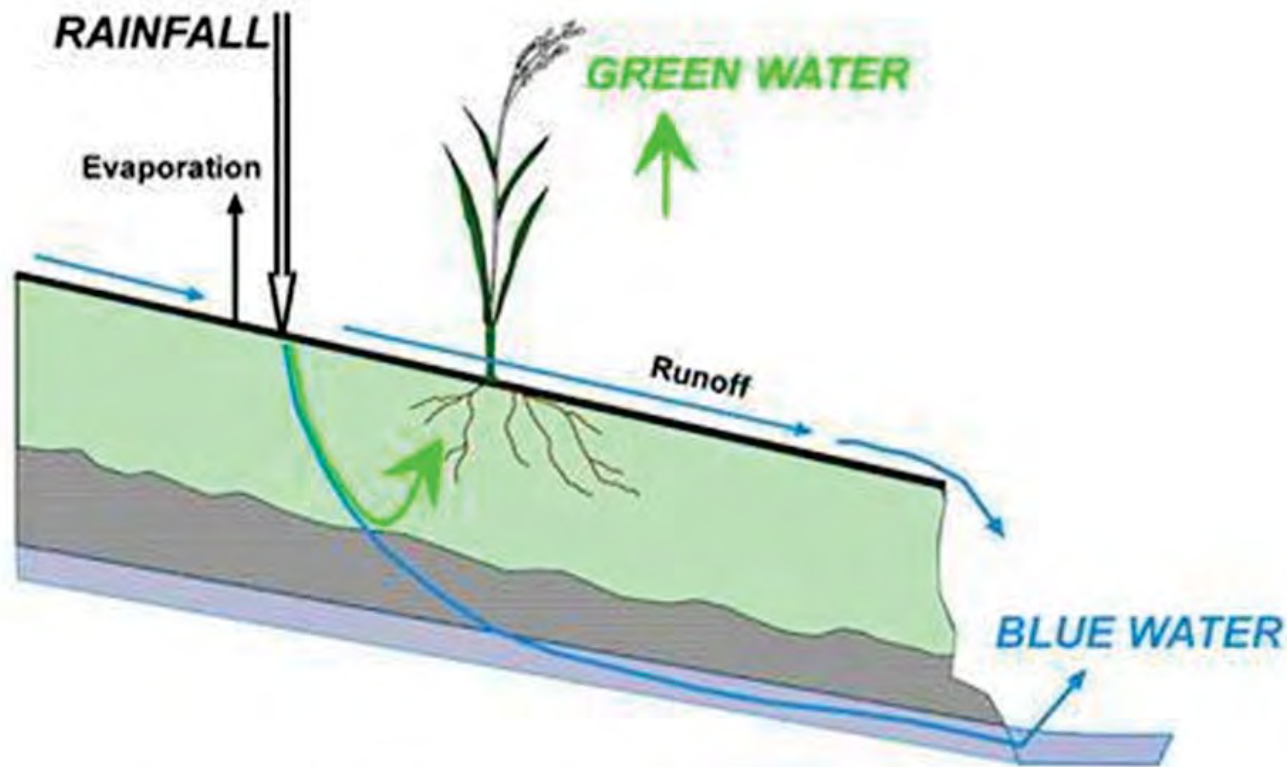
zai pits, Niger



infiltration trenches, Chile







Water, stored in the soil and used by plants, equals **green water**.
Runoff and deep drainage, recharging groundwater and feeding streams, equals **blue water**.

Rockström (1997)



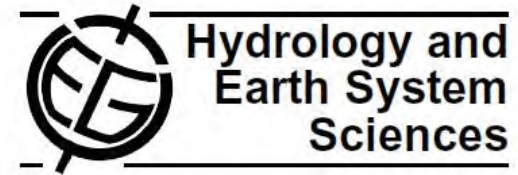
however... WHT often based on trial and error approach or at best on empirical approach

→ often unproductive, not efficient

→ impact on local/catchment hydrology unknown



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Modelling water-harvesting systems in the arid south of Tunisia using SWAT

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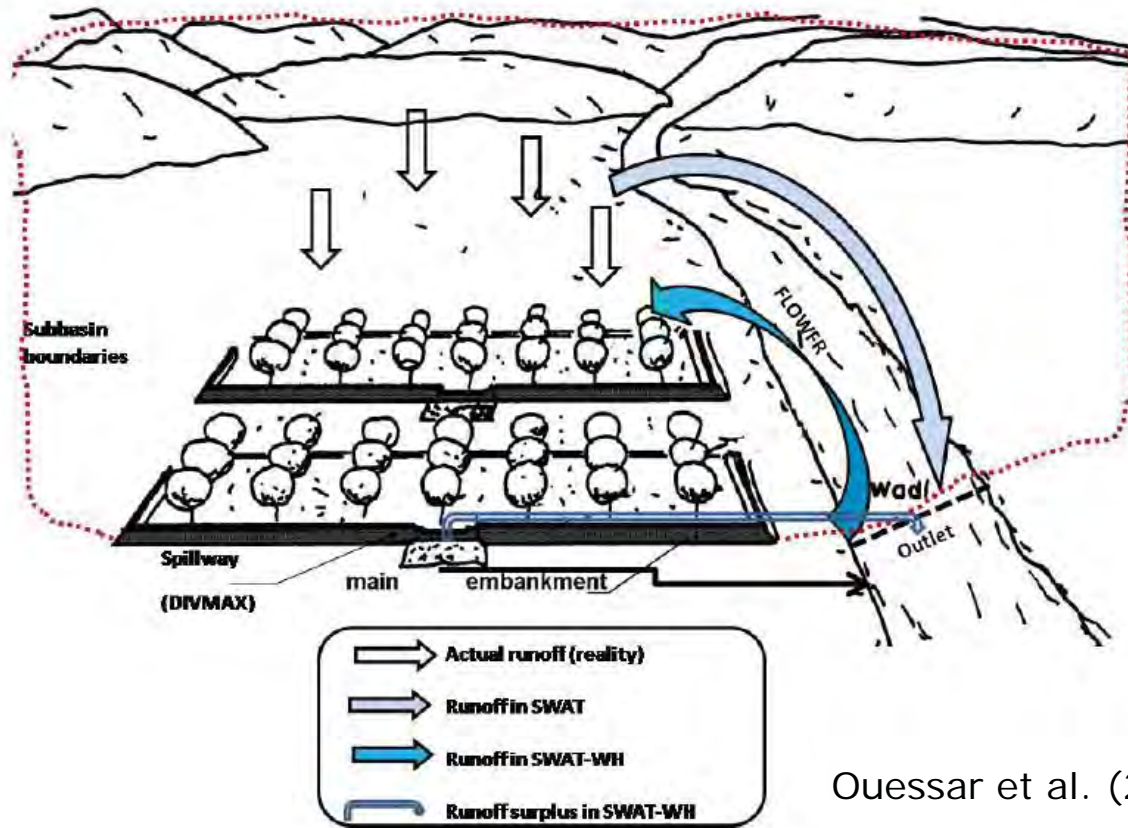
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SWAT

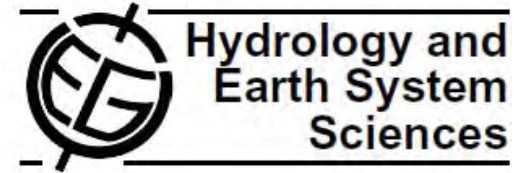


Ouessar et al. (2009, HESS)

but...not adapted for optimizing design



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Using an inverse modelling approach to evaluate the water retention in a simple water harvesting technique

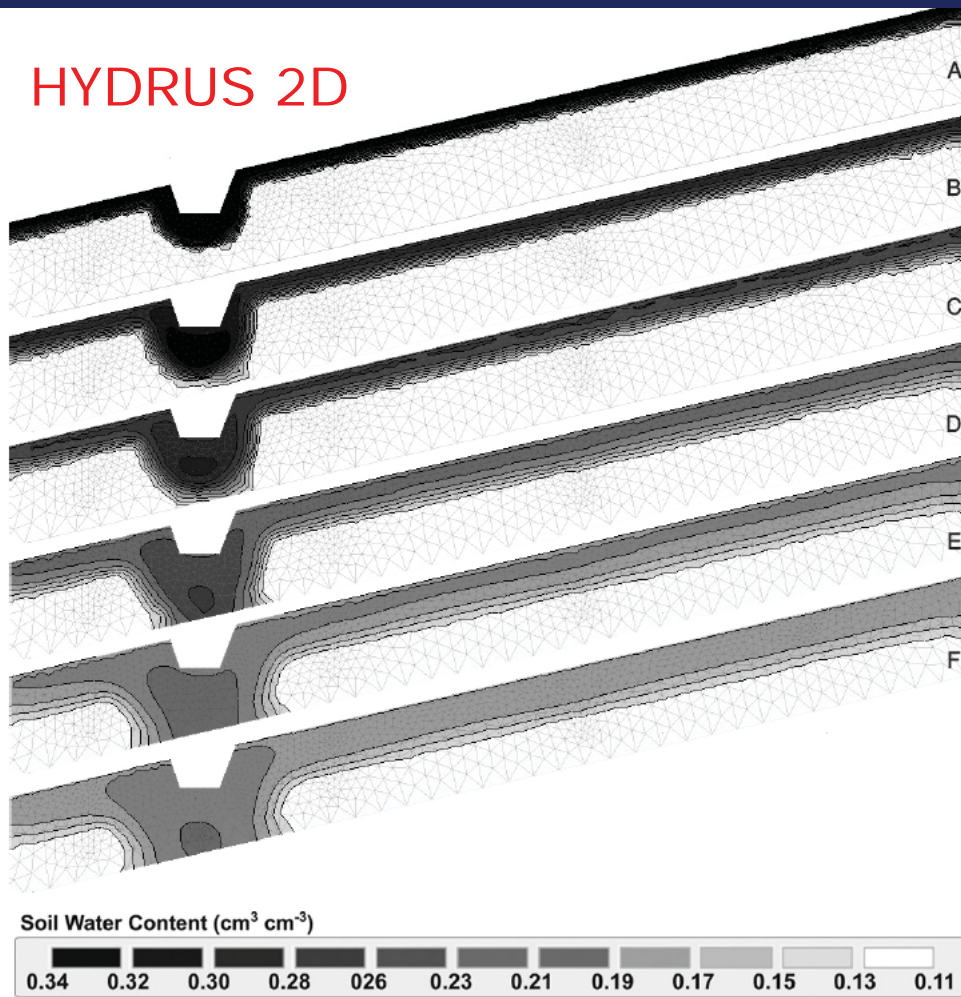
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HYDRUS 2D



Verbist et al. (2009, HESS)

but ... does not consider overland flow as such



→ fully coupled surface/subsurface flow model

based on blueprint by R.A. Freeze and R.L. Harlan.

“Blueprint for a physically-based digitally simulated, hydrological response model”. J. Hydrology 9:237-258, 1969

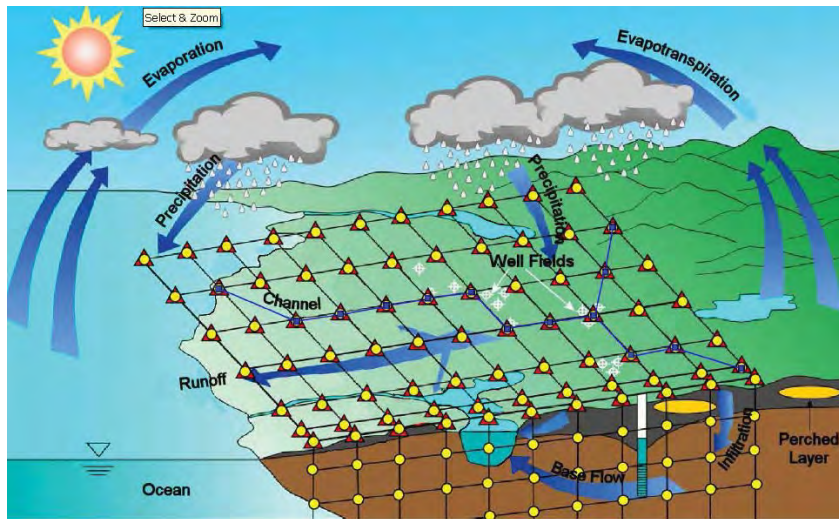
→ HydroGeoSphere (HGS)



- developed for “Simulating Flow and Contaminant Transport in Integrated Surface-Subsurface Flow Systems”
- at the University of Waterloo, Canada (Therrien, McClaren, Sudicky, Panday, 2009)
- first generation code: PhD of Joel VanDerKwaak, 1999, InHM (Integrated Hydrologic Model; U of Waterloo)



- Features:
 - Surface (overland) flow: 2D – Subsurface flow: 3D
 - Coupled Surface-Subsurface
 - Allows any time step
 - Allows any spatial resolution
 - Finite elements



Therrien et al. (2009)

Figure 1.2: Integrated Numerical Simulation of Hydrologic System.



- Governing equations:

- **porous medium:** 3D variably-saturated flow

Richards' equation:

$$-\nabla \cdot (w_m \mathbf{q}) + \sum \Gamma_{\text{ex}} \pm Q = w_m \frac{\partial}{\partial t} (\theta_s S_w)$$

Darcy-Buckingham equation:

$$\mathbf{q} = -\mathbf{K} \cdot k_r \nabla (\psi + z)$$

Mualem-van Genuchten

for S_w and k_r

w_m = porous medium vol. fraction

\mathbf{q} = fluid flux

Γ_{ex} = exchange flux

Q = source/sink

θ_s = sat. vol. water content

S_w = degree of saturation

\mathbf{K} = saturated hydraulic conductivity

k_r = relative hydraulic conductivity

ψ = pressure head

z = elevation



- Governing equations:

- **overland/stream:** 2D surface flow

diffuse wave approximation of Saint Venant equation:

$$-\nabla \cdot (d_o \mathbf{q}_o) - d_o \Gamma_o \pm Q_o = \frac{\partial \phi_o h_o}{\partial t}$$

flux equation:

$$\mathbf{q}_o = -\mathbf{K}_o \cdot k_{r_o} \nabla (d_o + z_o)$$

Manning equation:

$$K_{ox} = \frac{d_o^{2/3}}{n_x} \frac{1}{[\partial h_o / \partial s]^{1/2}}$$

d_o = water depth
 \mathbf{q}_o = fluid flux
 Γ_o = exchange flux
 Q_o = source/sink
 ϕ_o = surface flow porosity
 h_o = water surface elevation
 \mathbf{K}_o = surface conductance
 k_{r_o} = relative hydraulic cond.
 z_o = land surface elevation
 n_x = Manning roughness coeff.
 S = max. slope



- Governing equations:

– coupling?

subsurface 3D flow: $-\nabla \cdot (w_m \mathbf{q}) + \sum \Gamma_{ex} \pm Q = w_m \frac{\partial}{\partial t} (\theta_s S_w)$

surface 2D flow: $-\nabla \cdot (d_o \mathbf{q}_o) - d_o \Gamma_o \pm Q_o = \frac{\partial \phi_o h_o}{\partial t}$


→ first-order exchange:

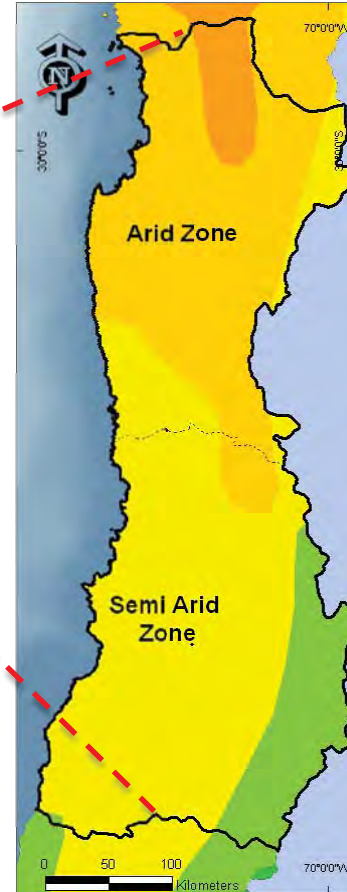
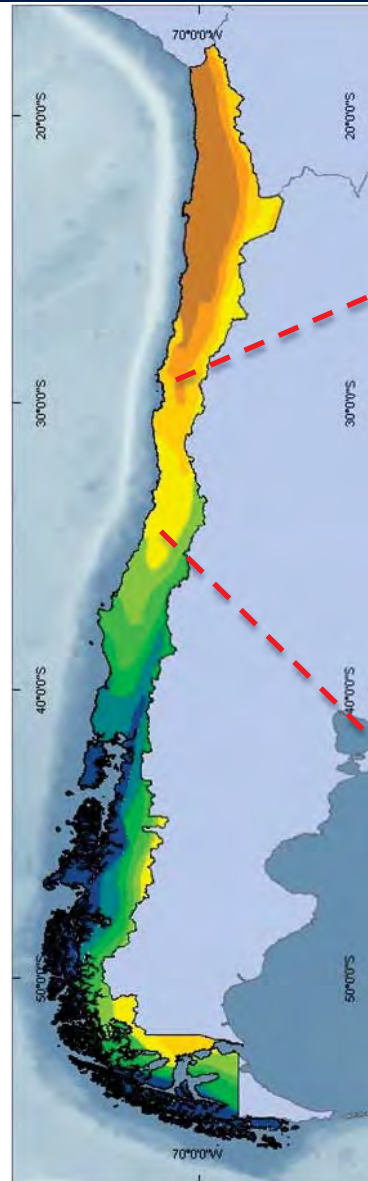
$$d_o \Gamma_o = \frac{k_r \mathbf{K}_{zz}}{l_{exch}} (h - h_o)$$

k_r = relative hydraulic conductivity of exchange flux
 \mathbf{K}_{zz} = sat. hydr. conduct. in vertical dir.
 l_{exch} = coupling length



3. Case studies

Aridity Regime	
	Xeric
	Hyper Arid
	Arid
	Semi Arid
	Subhumid
	Humid
	Hyper Humid
	Hydric
	Hyper Hydric



case study 1
small hillslope 6x2m

case study 2
small catchments ~3ha

Verbist et al. (2010)



(1) hillslope study

arid zone

$P = 99 \text{ mm}$

$ET_o = 1500 \text{ mm}$



(2) catchment study

semi-arid zone

$P = 560 \text{ mm}$

$ET_o = 1220 \text{ mm}$



in both cases: infiltration trenches were dug since '90s





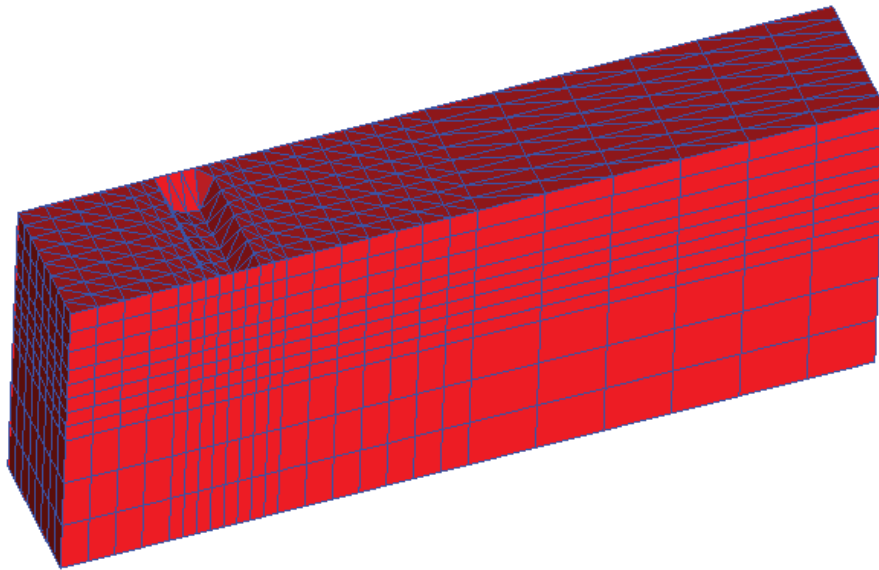
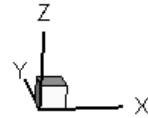
→ to combat desertification: by stimulating

- reforestation (*Eucalyptus*, *Pinus*)
- regeneration of natural vegetation (used by goats)



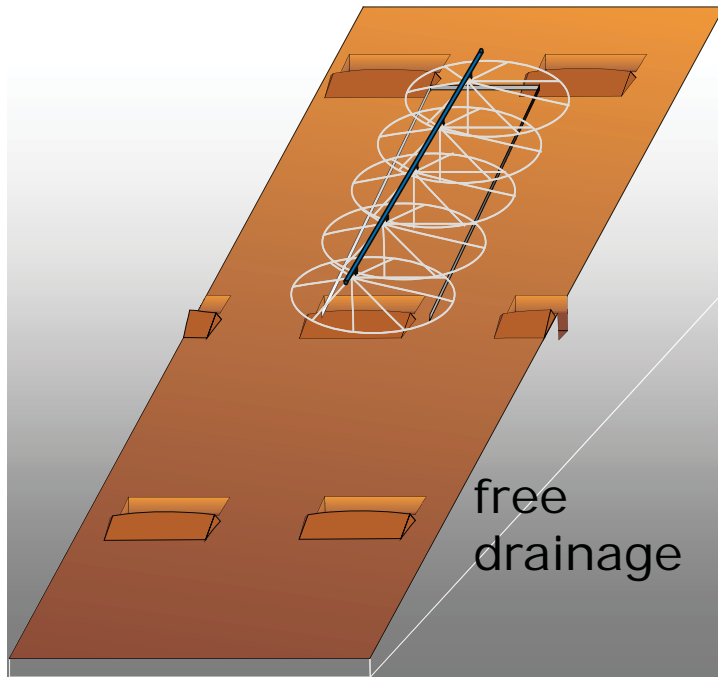
(1) Flow domain and boundary conditions:

- slope: 23%
- silt loam
- a field plot of 6 x 2 m
- one trench + catchment area
- 3D mesh



(1) Flow domain and boundary conditions (cont.):

- simulated rain (20 min, 120 mm h^{-1} , 7 nozzles)
+ evaporation (3.5 days)
- runoff and soil-water content measurements

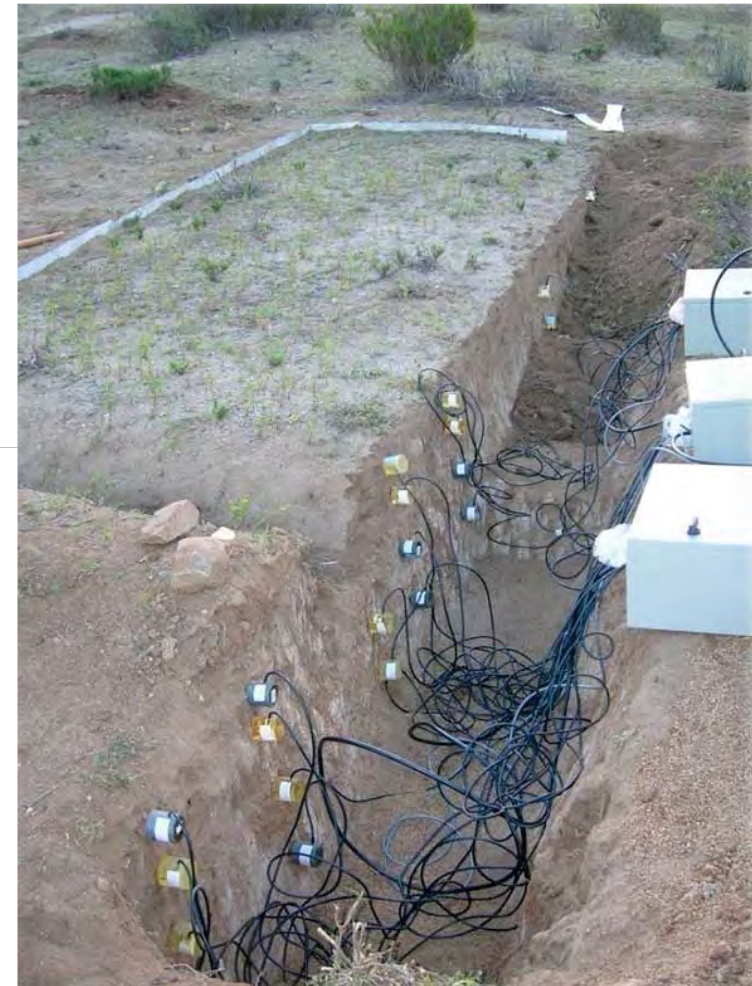
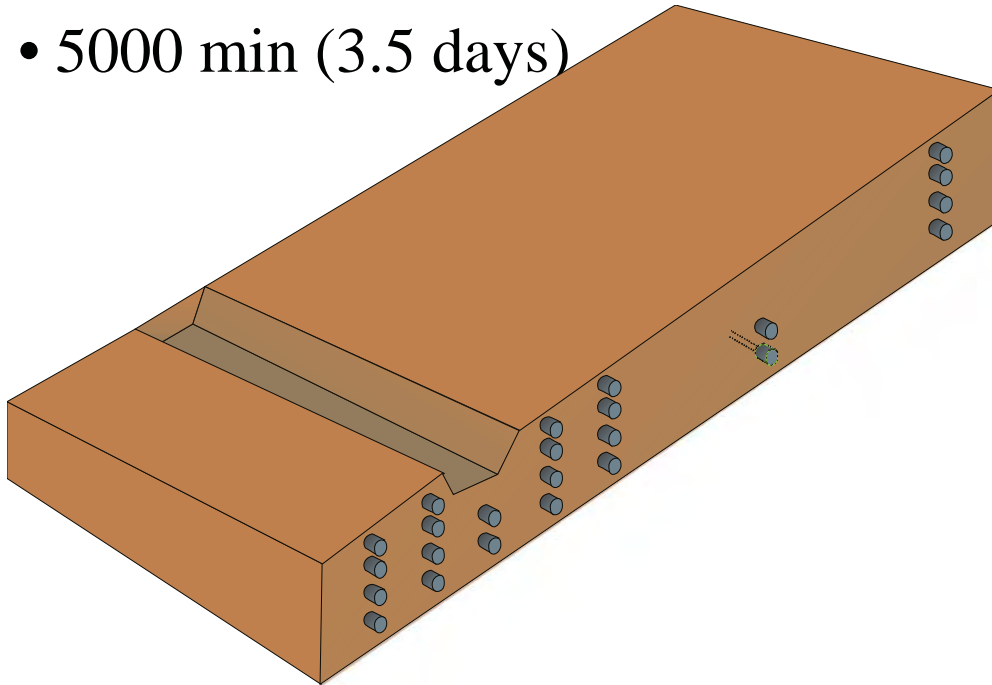


(2) model parameterisation: optimizing the runoff hydrograph
(from 10 independent rainfall simulations)



(2) model parameterisation (cont.): optimizing soil-water content time series

- 22 TDR probes
- probe length: 30 cm
- 5 min interval
- 5000 min (3.5 days)



(2) model parameterisation (cont.):

- 7 model parameters need to be estimated:

- K_{fs} : saturated hydraulic conductivity

- $\theta_r, \theta_s, \alpha, \beta, \lambda_p$: van Genuchten-Mualem WRC parameters

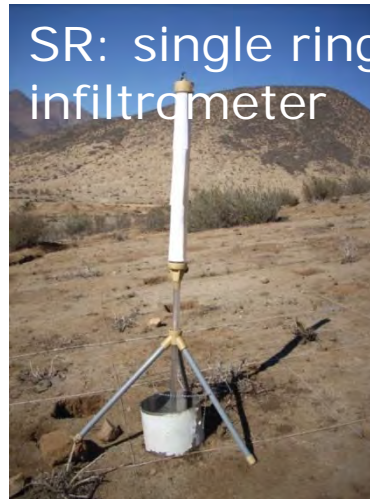
- n : Manning coefficient

- l_{exch} : coupling length (different in catchment area and reception area)

- for some, initial estimates from direct field/lab measurement



(2) model parameterisation (cont.): measurements of K_{fs} in 10 reps with 6 methods (side study)

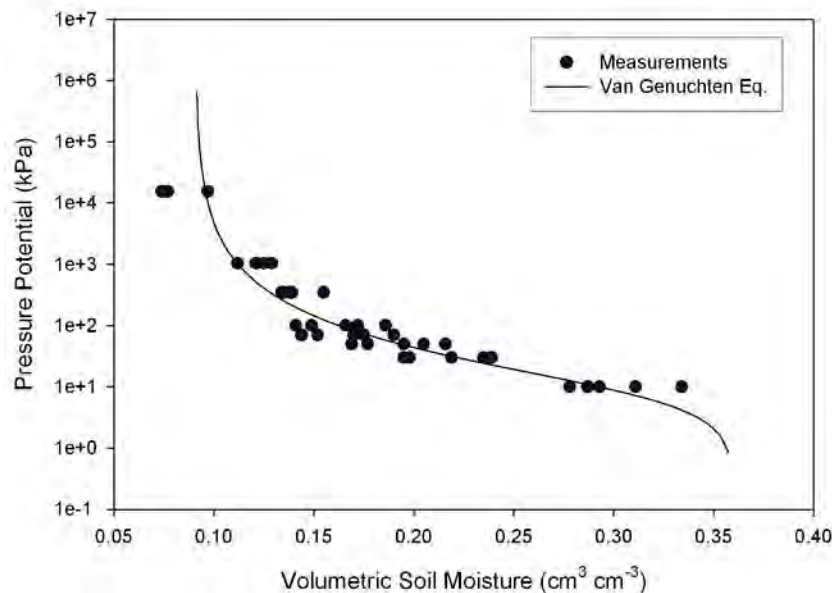


Verbist et al.
(2009, SSSAJ)
Baetens et al.
(2009, WRR)
Verbist et al.
(2010, VZJ)



(2) model parameterisation (cont.): water retention curves

- undisturbed soil cores (Kopecky – 100 cm³)
- tension table (Eijkelkamp Agr. Eq.)
0-10 kPa
- pressure chambers (Soilmoisture Eq.)
20-1500 kPa



(2) model parameterisation (cont.):

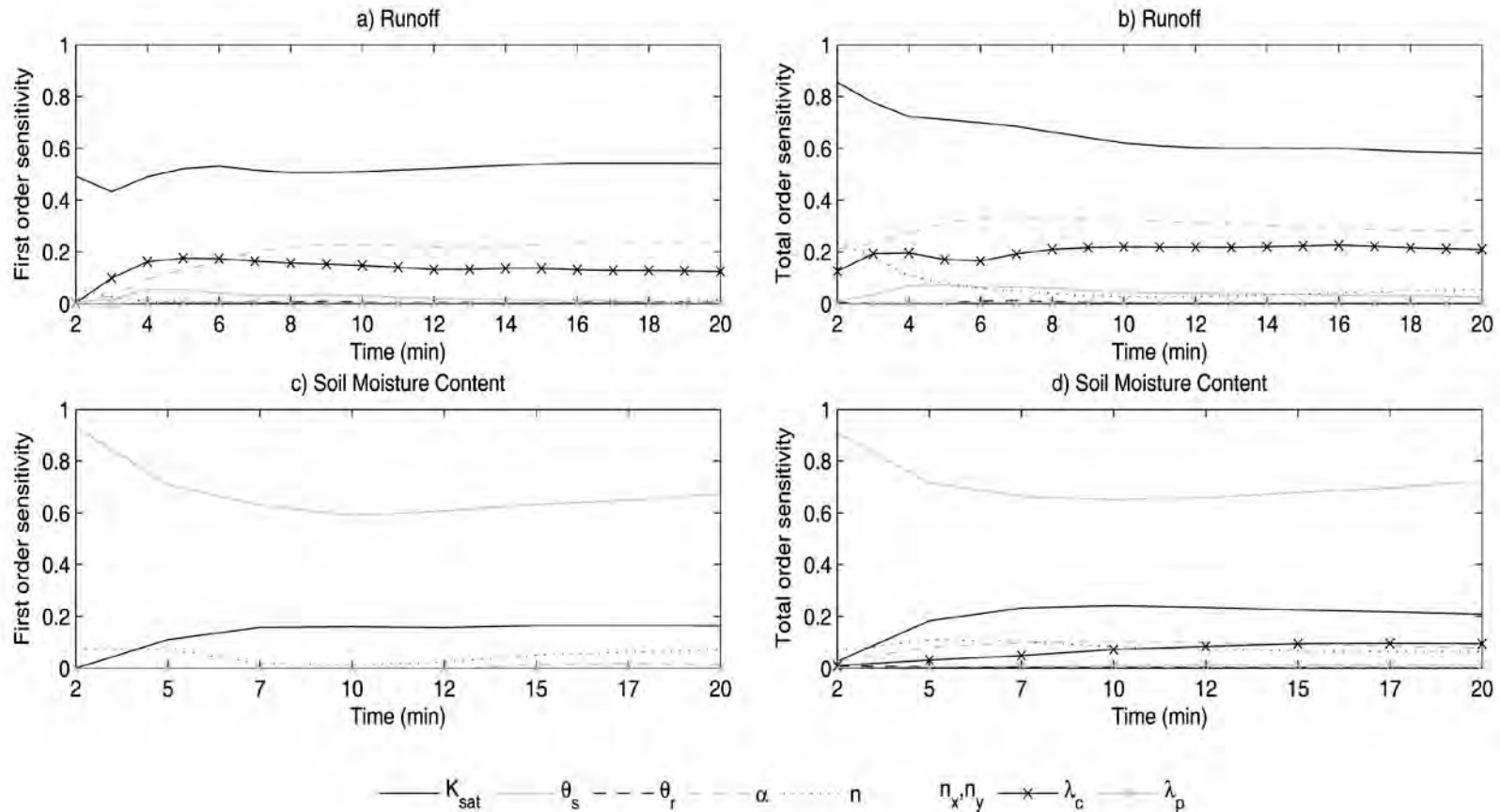
variance-based global sensitivity analysis (GSA) using Jansens' estimator which applies a quasi-random sequence generator

→ 5000 model runs using the parallel HGS version (Park et al., 2010) on 16 computer cores (Intel Xeon L5420 2.5 GHz)

(no local sensitivity analysis or one-at-a-time approach since in non-linear models, parameter interaction might occur)



(2) model parameterisation (cont.):



→ insensitive for n_x, n_y, λ_p → fixed

Verbist et al. (2012, VZJ)



(2) model parameterisation (cont.): parameter estimation

objective function: minimisation-algorithm of **Levenberg–Marquardt** (1963)

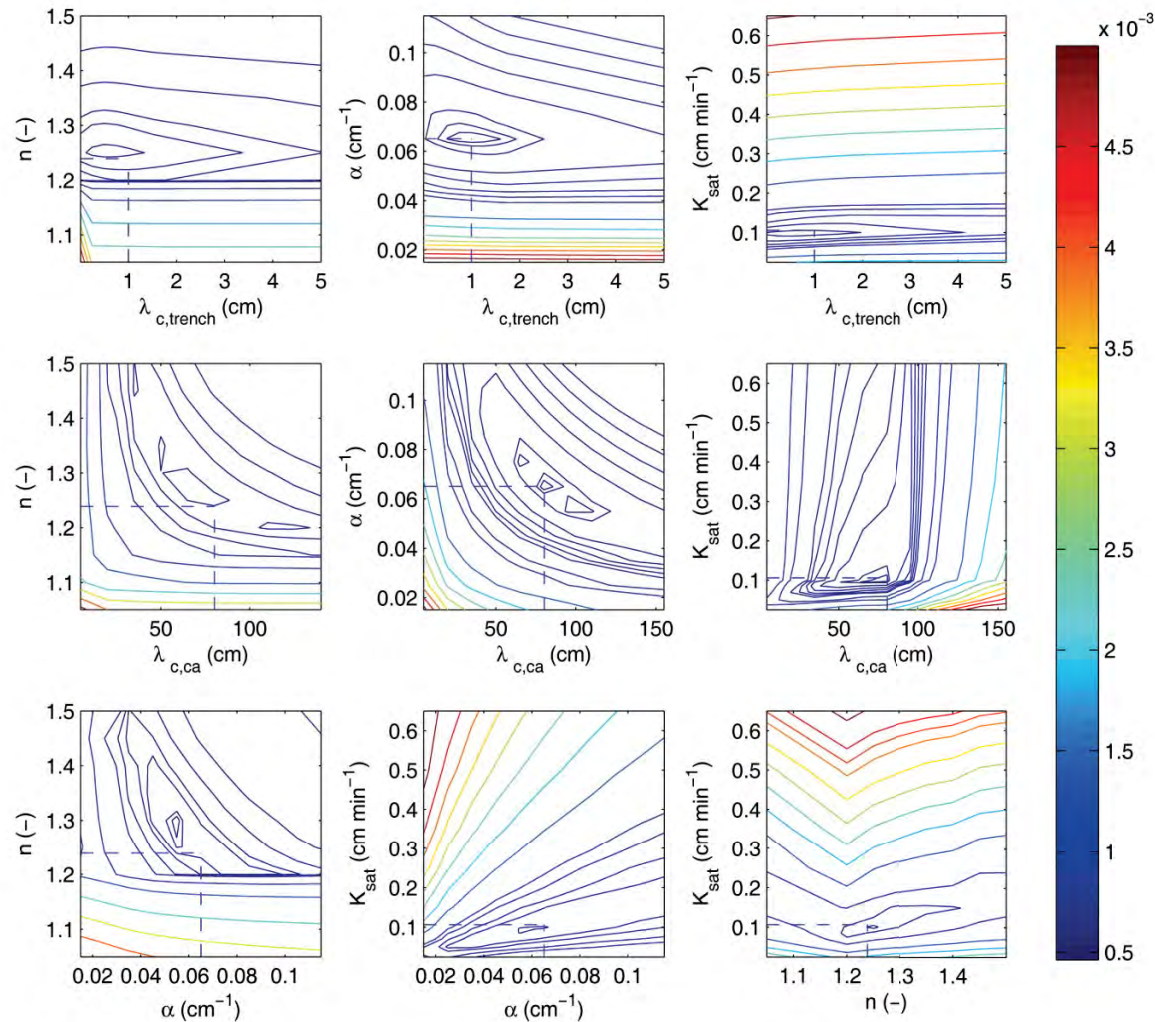
$$\Phi[\Theta(t), R(t)] = \sum_{j=1}^m \sum_{i=1}^{n_1} u_i [\theta_o(x_j, t_i) - \theta_s(x_j, t_i, \beta)]^2 + \sum_{i=1}^{n_2} v_i [R_o(t_i) - R_s(t_i, \beta)]^2$$

Coupling HydroGeoSphere with **PEST** (**Parameter Estimation Software**) (Doherty, 2010)

(**DREAM**, **DiffeRential Evolution Adaptive Metropolis**)
(Laloy & Vrugt, 2012)



3a. Hillslope study

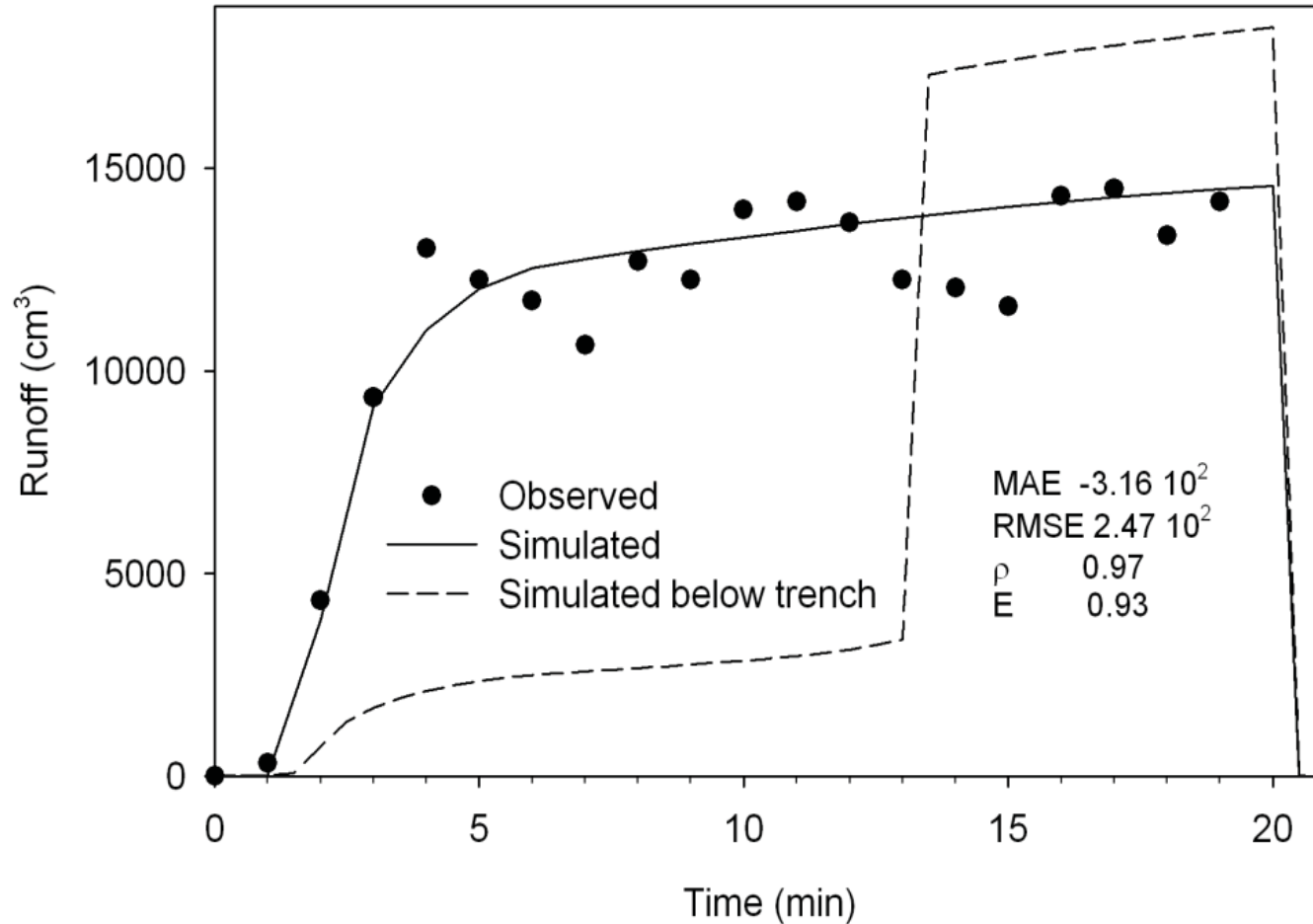


Verbist et al.
(2012, VZJ)

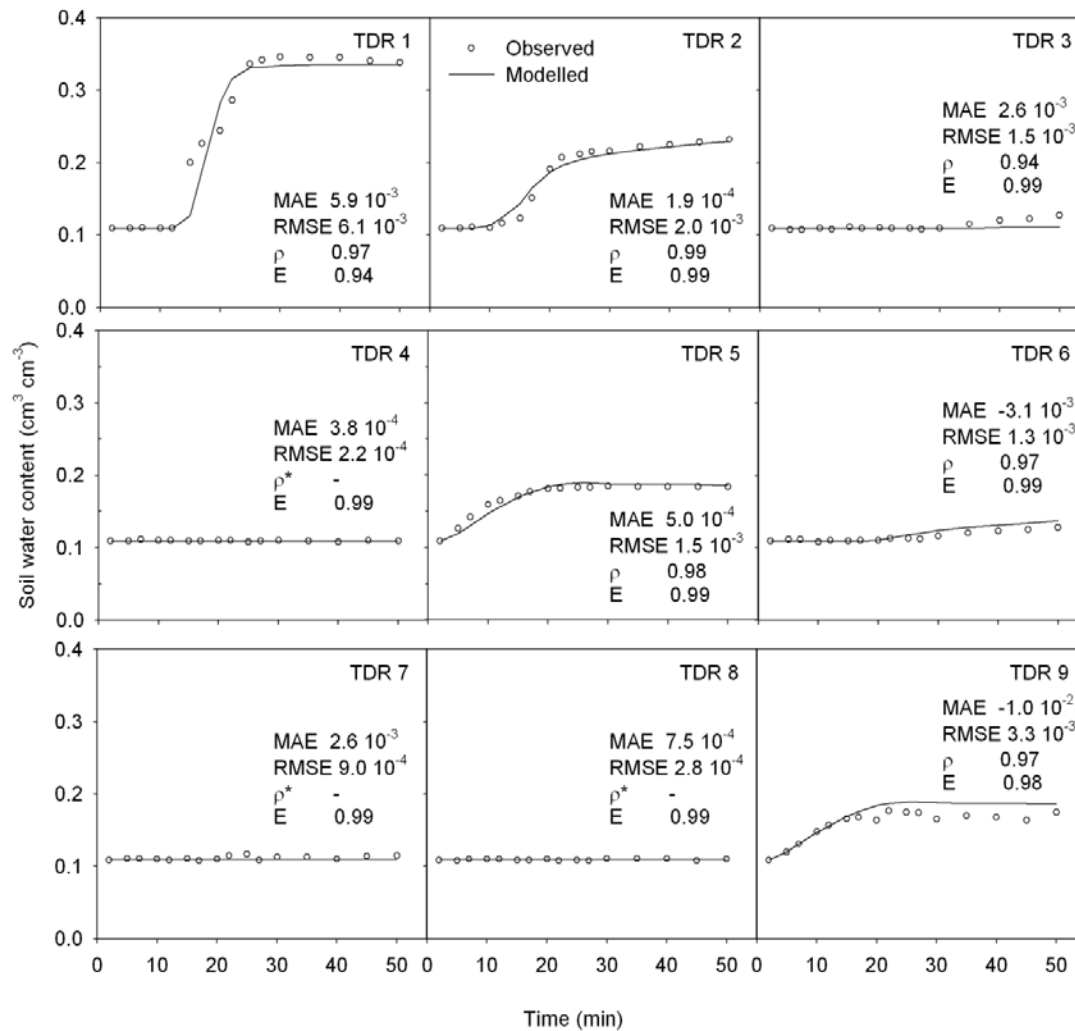
→ well-posed model with unique inverse solution



(3) model results: simulation of runoff

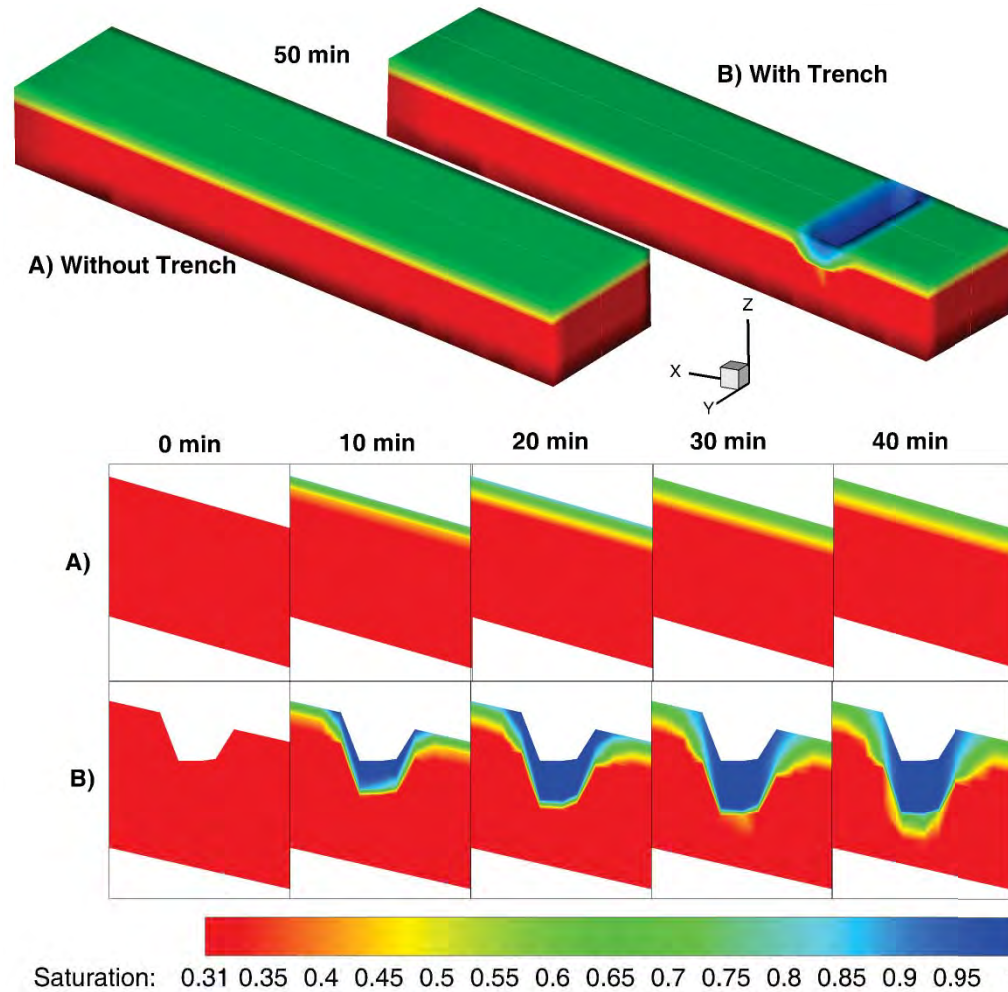


Verbist et al. (2012, VZJ)

(3) model results (cont.): simulation of **water content** with time

Verbist et al.
(2012, VZJ)

(3) model results (cont.) : trench filling + redistribution

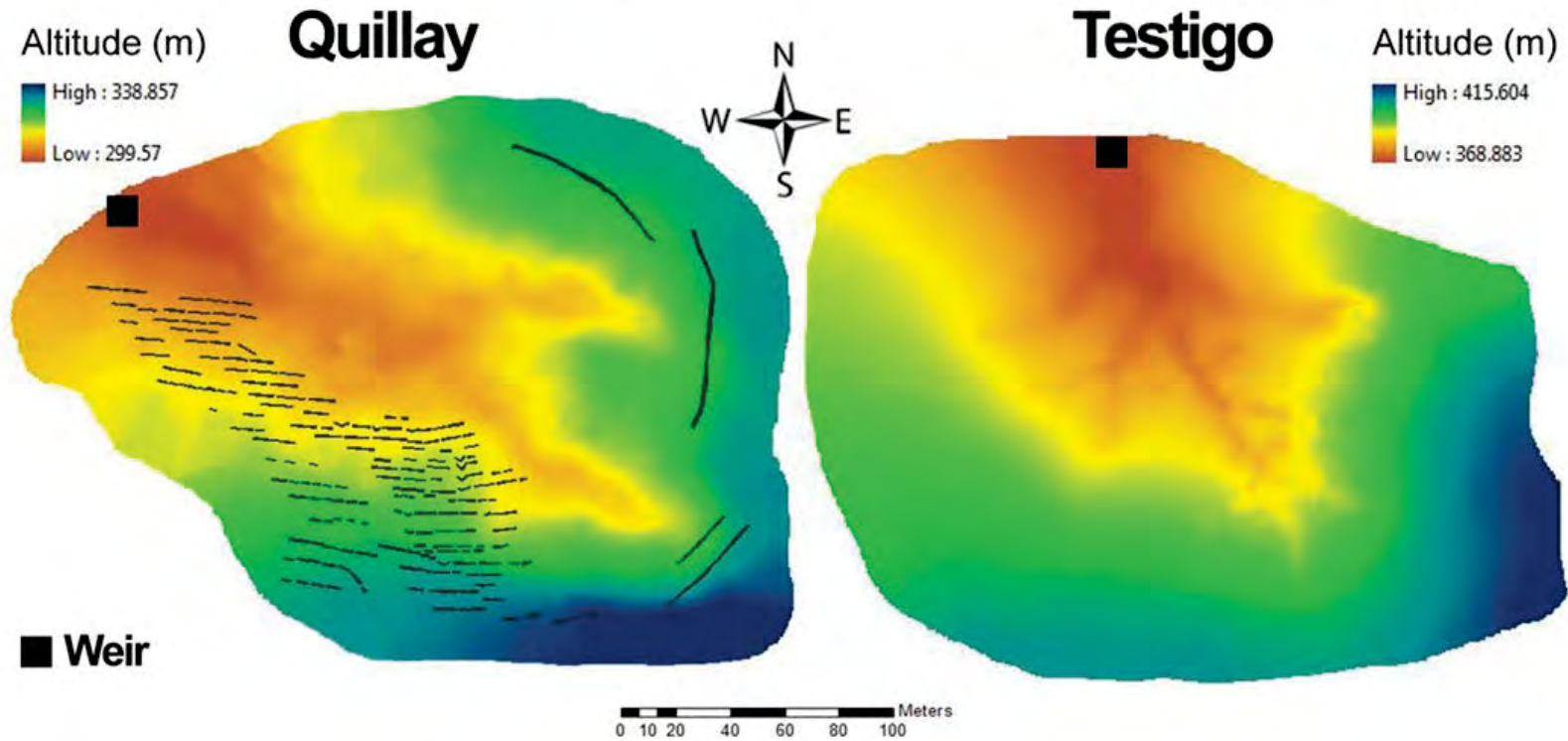


Verbist et al. (2012, VZJ)

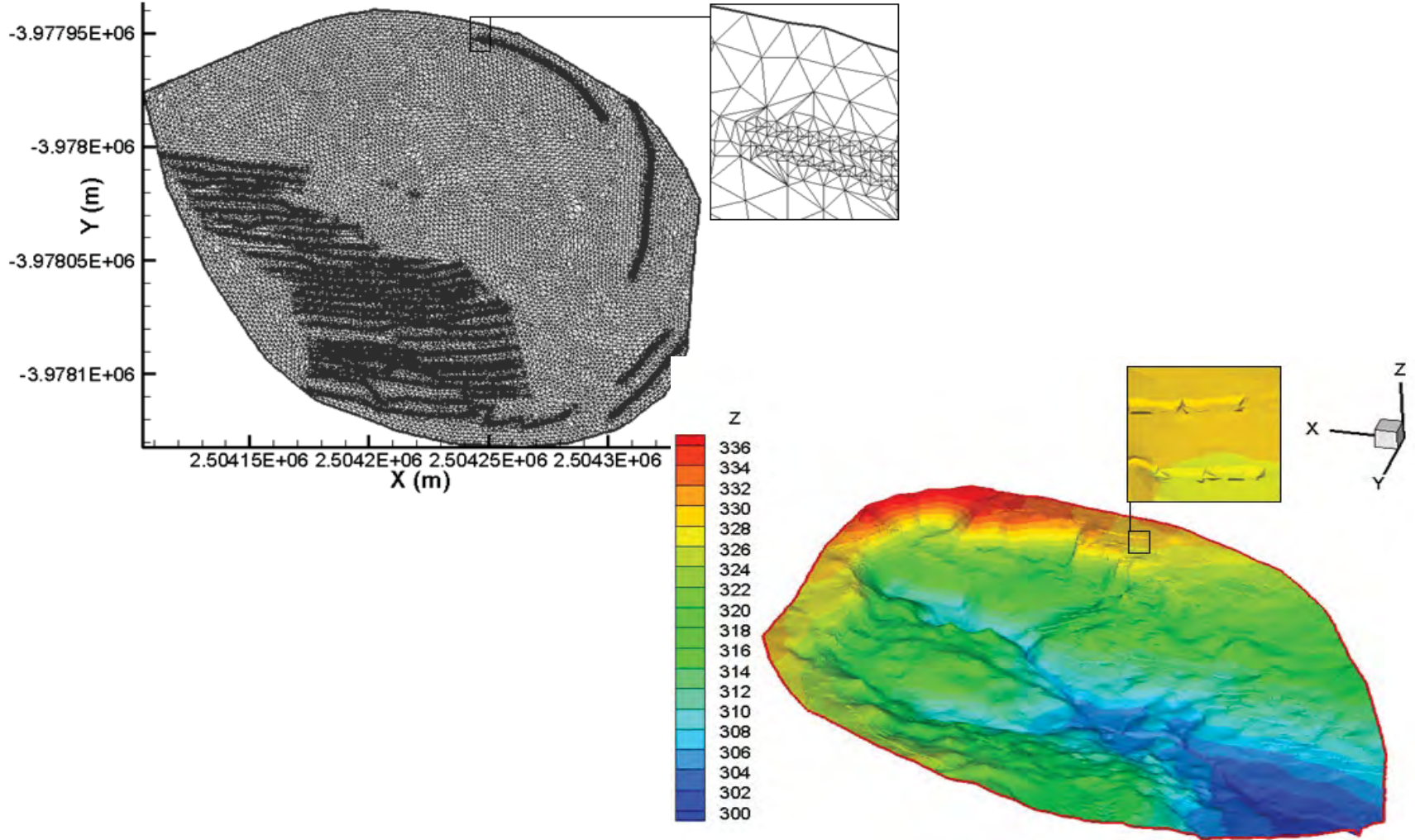


(1) flow domain + boundary conditions:

- ~3 ha
- loamy soils
- natural conditions



(1) flow domain + boundary conditions:

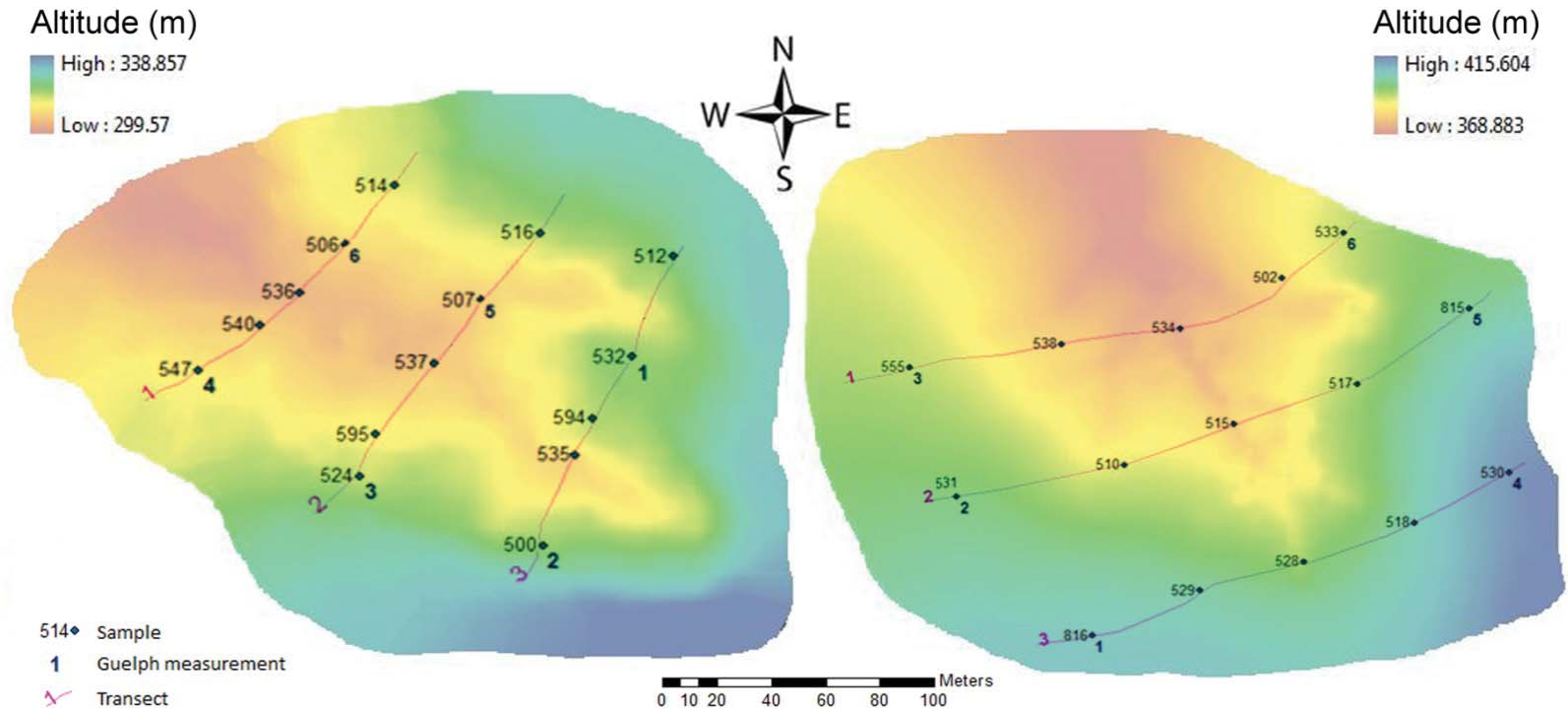


(2) model parameterisation:
rainfall simulations 1x1 m plots (runoff + TDR water content)



(2) model parameterisation (cont.):

K_{fs} and WRC at 5 locations along 3 transects (9 reps per loc.)



- (2) model parameterisation (cont.):
validation with runoff discharge data from outlet of watershed



(3) model results: simulation of pressure head at DOY1997

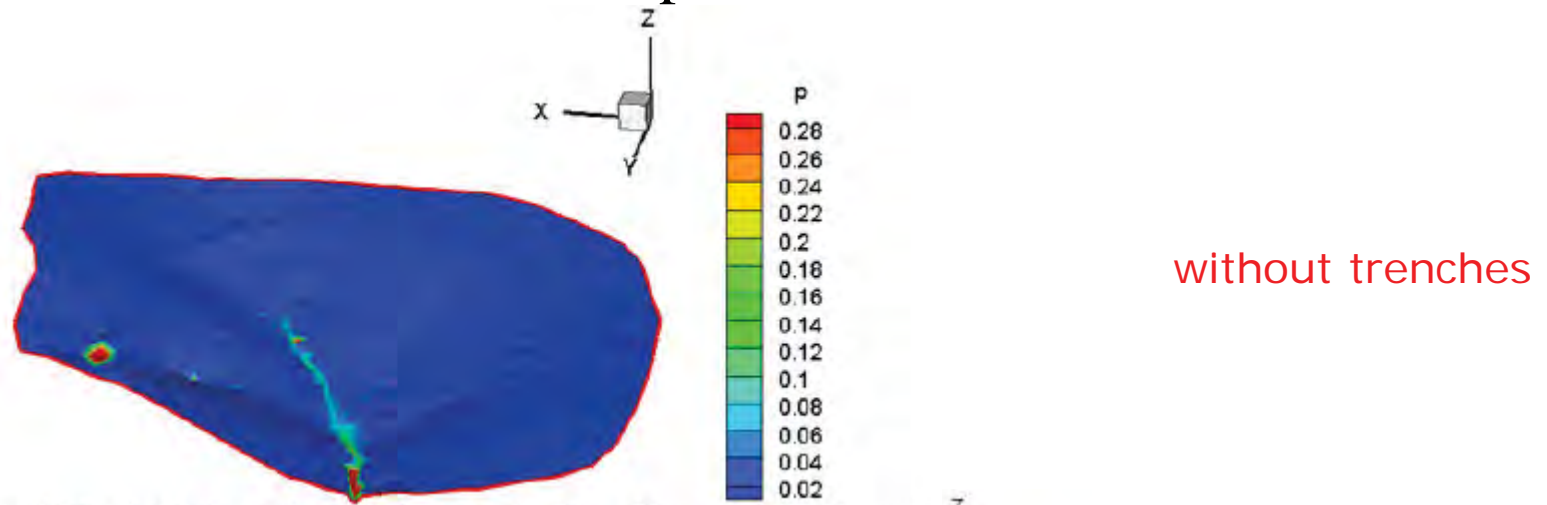


Figure 4.13: Tecplot 3D-visualization of Testigo with a simulation of the pressure head (meters)

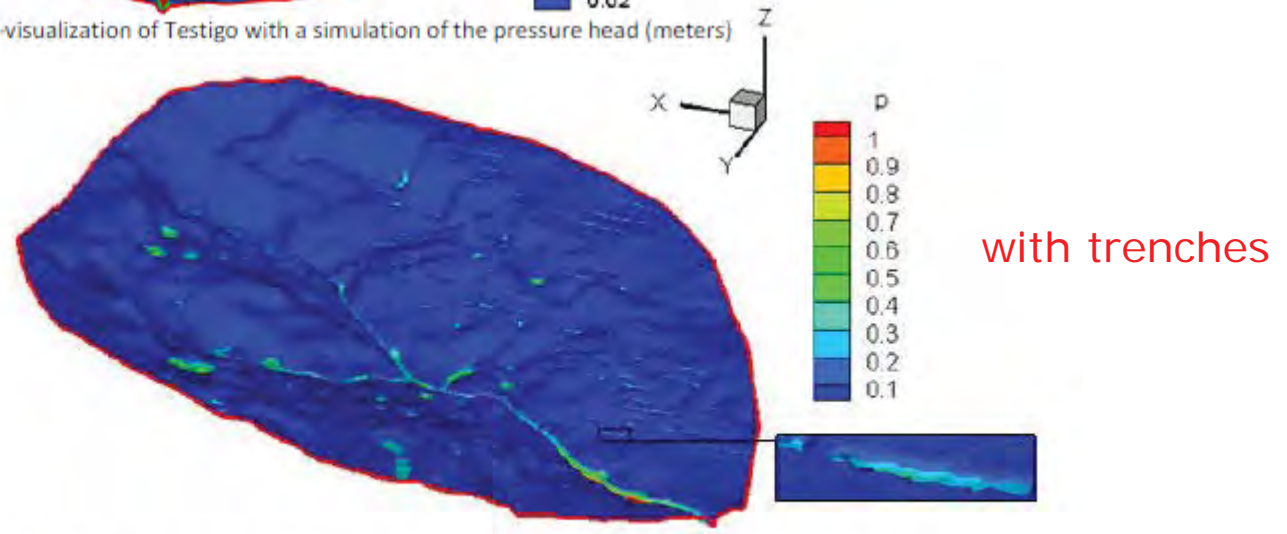
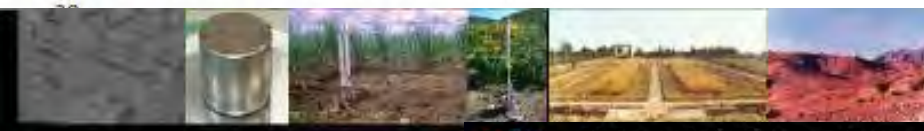
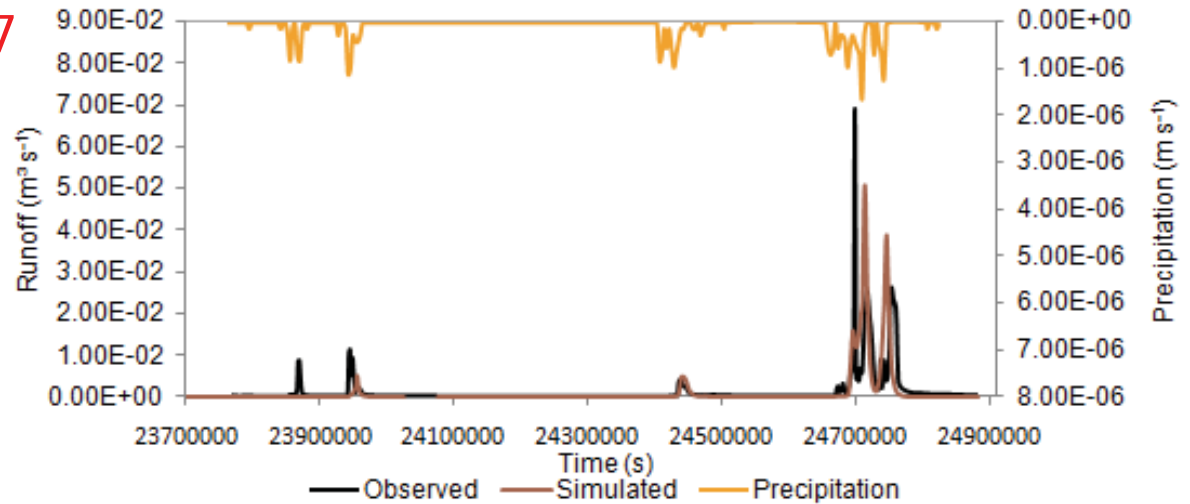


Figure 4.14: Tecplot 3D-visualization of Quillay with a simulation of the pressure head (meters)

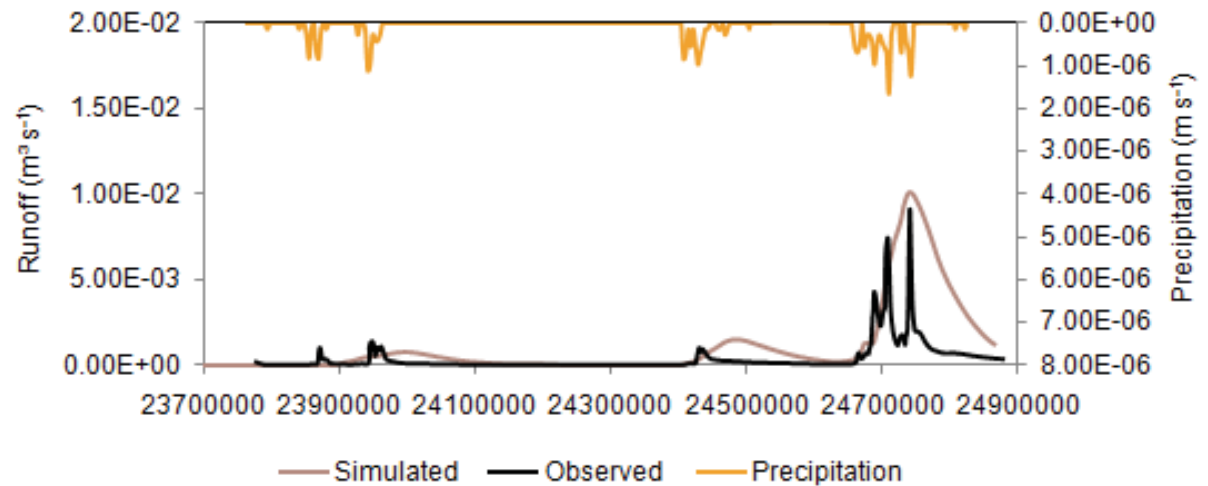


(3) model results: simulation of runoff discharge

without WHT 1997



with WHT 1997



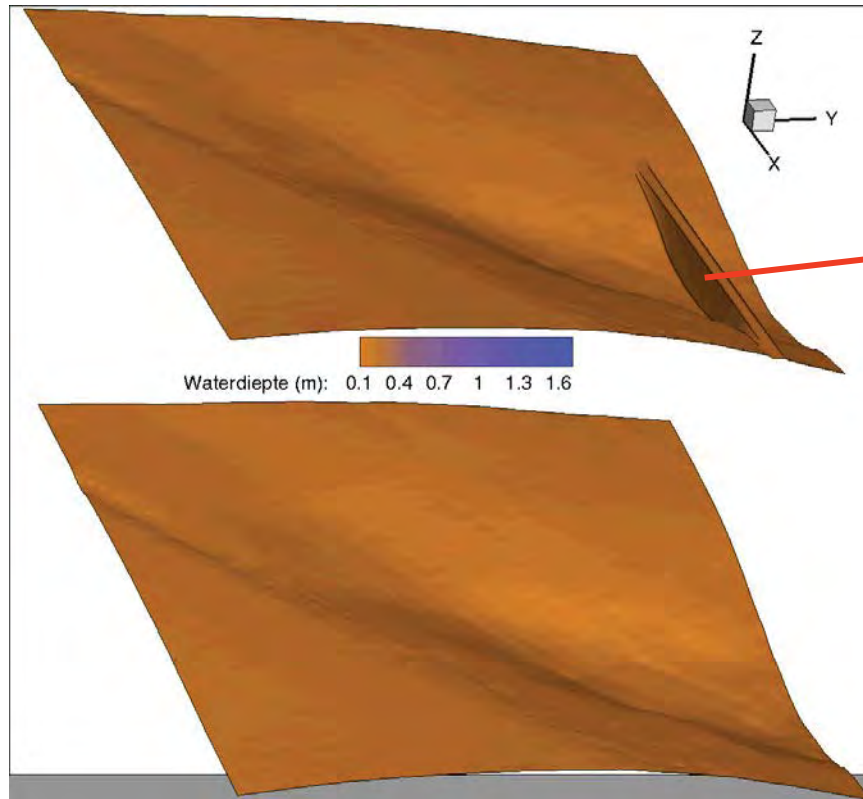
- model most sensitive to K_{fs} , van Genuchten parameters and coupling length
- enables to mimic runoff and changes in SWC in small hillslope plots and runoff discharge from watersheds → water balance
- versatile tool for improving design of WHT and evaluating their impact for current, past and future climates
- visualisation of water flow
- allows to evaluate upstream and downstream impact of WHT at watershed scale



- future work: testing for other techniques



- future work (cont.): sediment transport model



Thank you!

